

Wind fields in the Martian atmospheric boundary layer obtained by high-resolution large eddy simulations

*Kuriki Murahashi¹,

Kohei Suto¹, Seiya Nishizawa², Masaki Ishiwatari¹,
Masatsugu Odaka¹, Kensuke Nakajima³, Shin-ichi Takehiro⁴,
Ko-ichiro Sugiyama⁵, Hiroataka Ogihara¹,
Yoshiyuki O. Takahashi⁶, Yoshi-Yuki Hayashi⁶

1. Hokkaido Univ., 2. RIKEN/AICS, 3. Kyushu Univ.,
4. Kyoto Univ., 5. National Institute of Tech. Matsue college,
6. Kobe Univ.

Self introduction

- Doctor course student
 - This presentation shows planning of my doctor thesis
- Sorry for be late, because of business trip to Nayoro observatory



- Dormitory leader of Keiteki-ryo dormitory (2013 - 2014)

It's me!



One third of Keiteki-ryo Dormitory residents (2014)

Outline

■ Introduction

- Dust in the Martian atmosphere
- MGCM calculation including dust processes
- Problems of dust lifting schemes used in MGCM
- Purpose of this study

■ LES Model / Data

■ Results

- Wind fields of the highest resolution results
- Surface wind stress of various resolution results
- Wind structures associated with the strongest wind stress
- Dust flux distribution

■ Summary

Dust in the Martian atmosphere

- Dust in the Martian atmosphere greatly influences optical depth and temperature structure. (Smith, 2009, etc.)
- Various space-time scale dust phenomena exist.



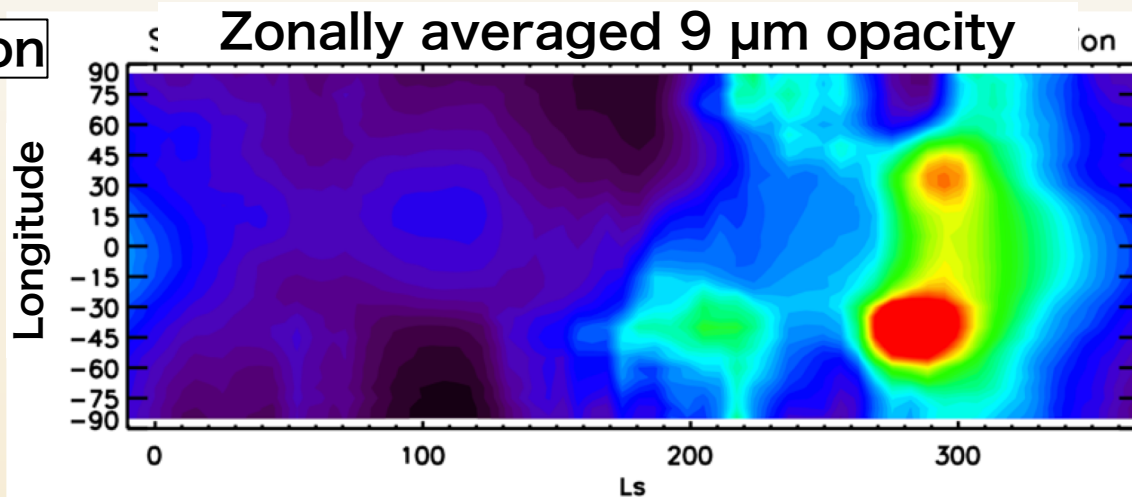
		Global Dust Storm	
Months		Regional Dust Storm	
Days		Local Dust Storm	
Minutes	Dust Devil		
Turbulence			
	10 m	10 km	10,000 km

Dust Devils (Recorded by Rover Spirits)
<http://mars.nasa.gov/mer/gallery/press/spirit/20050819a.html>

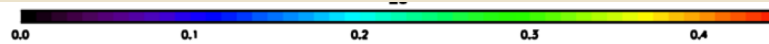
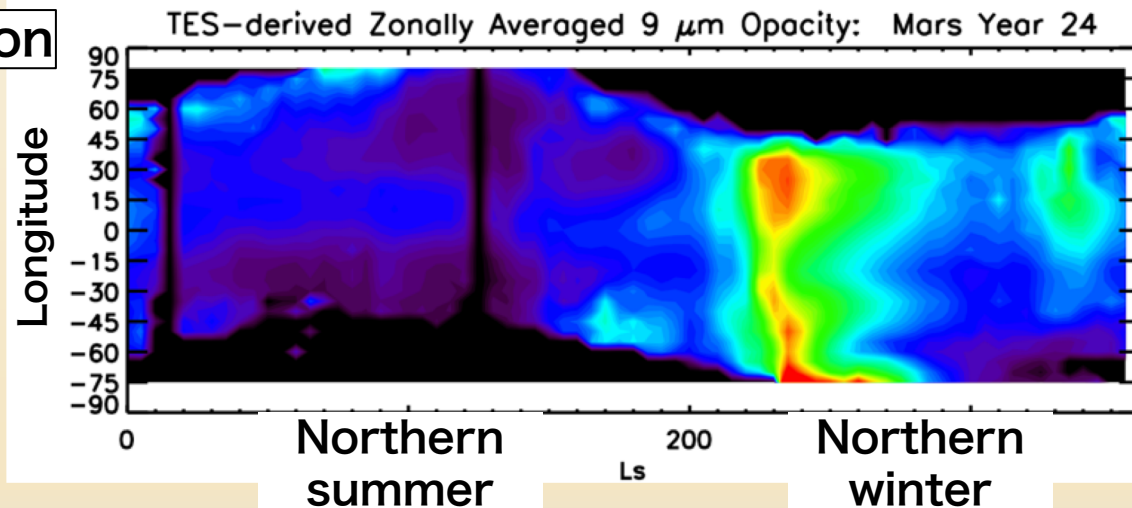
MGCM calculation including dust processes

- Kahre et al. (2006) simulates seasonal variability of dust distribution.

Simulation

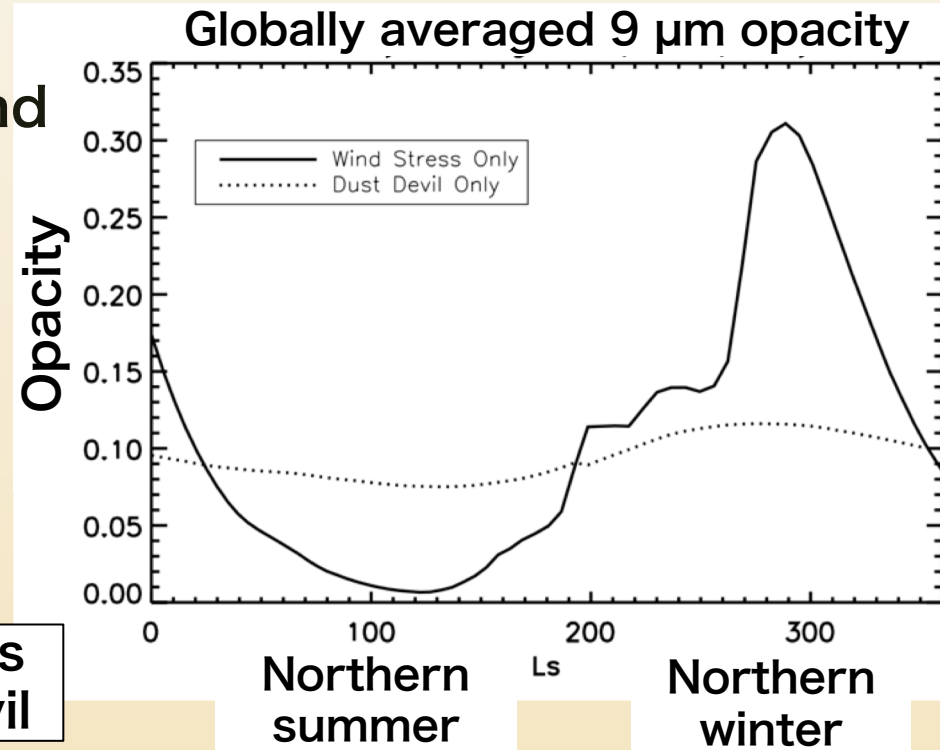


Observation



Dust lifting schemes used in MGCM

- Kahre et al. (2006) uses 2 types of dust lifting parameterization schemes.
 - Wind stress lifting schemes
 - The seasonal variability of dust amount can be simulated.
 - Dust devil lifting schemes
 - An amount of background dust can be simulated.



Solid line : wind stress
Dotted line : dust devil

Wind stress lifting schemes

■ KMH scheme (Kahre et al., 2006)

$$F_W = \alpha_W \times 2.3 \times 10^{-3} \tau^2 \left(\frac{\tau - \tau^*}{\tau^*} \right)$$

Parameters in Kahre et al. (2006)

F_W : Dust flux [kg/(m² s)]

α_W : Efficiency factor

τ : Surface wind stress [N/m²]

τ^* : Threshold value [N/m²]

τ^*	α_W
10×10^{-3}	0.02
22.5×10^{-3}	0.1
35×10^{-3}	0.45

■ Based on observational results on the Earth (Sahara desert). (Westphal et al., 1987)

■ Adjusting to the Martian conditions. (Kahre et al., 2006)

- Atmospheric density, gravitational acceleration.

Dust devil lifting schemes

- DDA scheme (Newman et al., 2002)

$$F_D = \alpha_D F_s (1 - b) \quad b = \frac{p_s^{\chi+1} - p_{con}^{\chi+1}}{(p_s - p_{con})(\chi + 1)p_s^\chi} \quad \chi \equiv \frac{R}{c_p}$$

F_D : Dust flux [kg/(m² s)]

F_s : Sensible heat flux [W/m²]

α_D : Efficiency factor [kg/J]

p_s : Surface Pressure [Pa]

p_{con} : Pressure at the top of PBL [Pa]

R : Specific gas constant

c_p : Specific heat capacity

- Based on the thermodynamics of dust devils as a heat engine.
(Rennò et al., 1998)
- Thermal efficiency is used for expressing dust flux.
 - With the higher sensible heat flux, kinetic energy of convection becomes larger.
 - With the higher PBL altitude, the conversion rate of kinetic energy from sensible heat flux becomes larger.
 - Therefore the amount of dust flux increases.

Problems of parameterization

- Adjusting parameters are necessary in order to simulate observational results.
 - Wind stress threshold value should be decreased compared to experimental value. (Greeley and Iversen, 1985)
- Schemes have been developed without considering details of wind structures.
- Wind stress schemes are suspected to include effects of dust devil schemes.

Purpose of this study

- Our purpose is to reconsider the schemes with examining relationship between wind microstructures such as dust devils and large scale convective structures.
 - What are characteristics of the wind field?
 - How much is the strength of the wind stress?
- In this study, we examine LES with several km domain.
 - Our results can be applied to MGCM.
 - The most high-resolution MGCM can resolved up to several km. (~ 11 km; Takahashi et al., 2011)

The highest resolution LES for the Mars

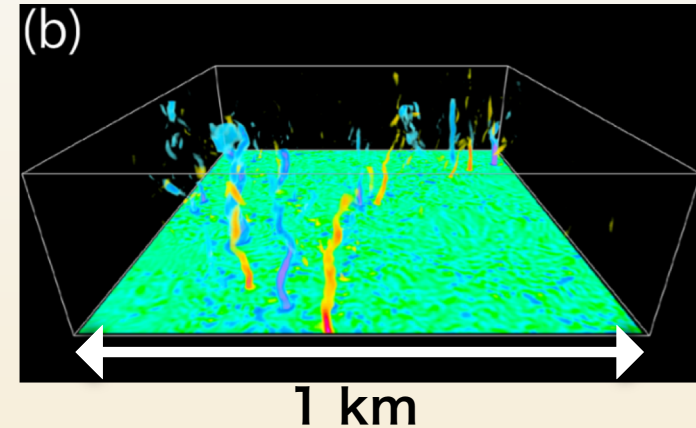
■ Nishizawa et al. (2016)

- Domain :
Horizontal 19.2 km, Vertical 21 km
- Horizontally periodic boundary conditions.
- Resolution 5, 10, 25, 50, 100 m
 - About 4.8×10^{10} grid points in 5 m resolution.
(1 time snapshot has 1.2 TB !)

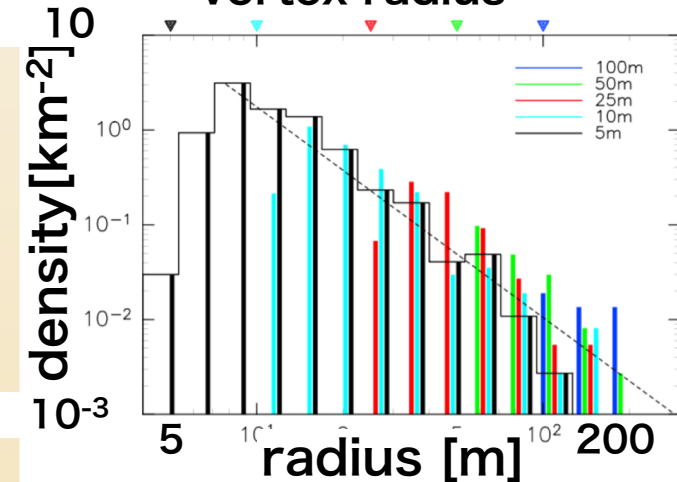
■ Statistics on vortices are investigated.

- e.g. vortex radius distribution of 62.5 m height at LT = 14:30

Vorticity distribution



Frequency distribution of vortex radius



Nishizawa et al. (2016)

Model

- **SCALE-LES** (Nishizawa et al., 2015; Sato et al., 2015)
 - <https://scale.aics.riken.jp/index.html>
 - 3D fully compressible non-hydrostatic equations model.
 - Developed by RIKEN/AICS.
- **Turbulence process**
 - Smagorinsky-type eddy viscosity model (Brown et al., 1994; Scotti et al., 1993).
- **Surface model**
 - Louis-type bulk method (Louis 1979, Uno et al., 1995).

Settings

■ Thermal forcing

- The heating rate and surface temperature are given by one-dimensional simulation by Odaka et al. (2001)

■ Initial State

- 10 - 100 m resolution :
Stable stratified stationary atmosphere with tiny random temperature perturbations.
- 5 m resolution :
interpolated 10 m result on LT = 14:00

■ Integration period

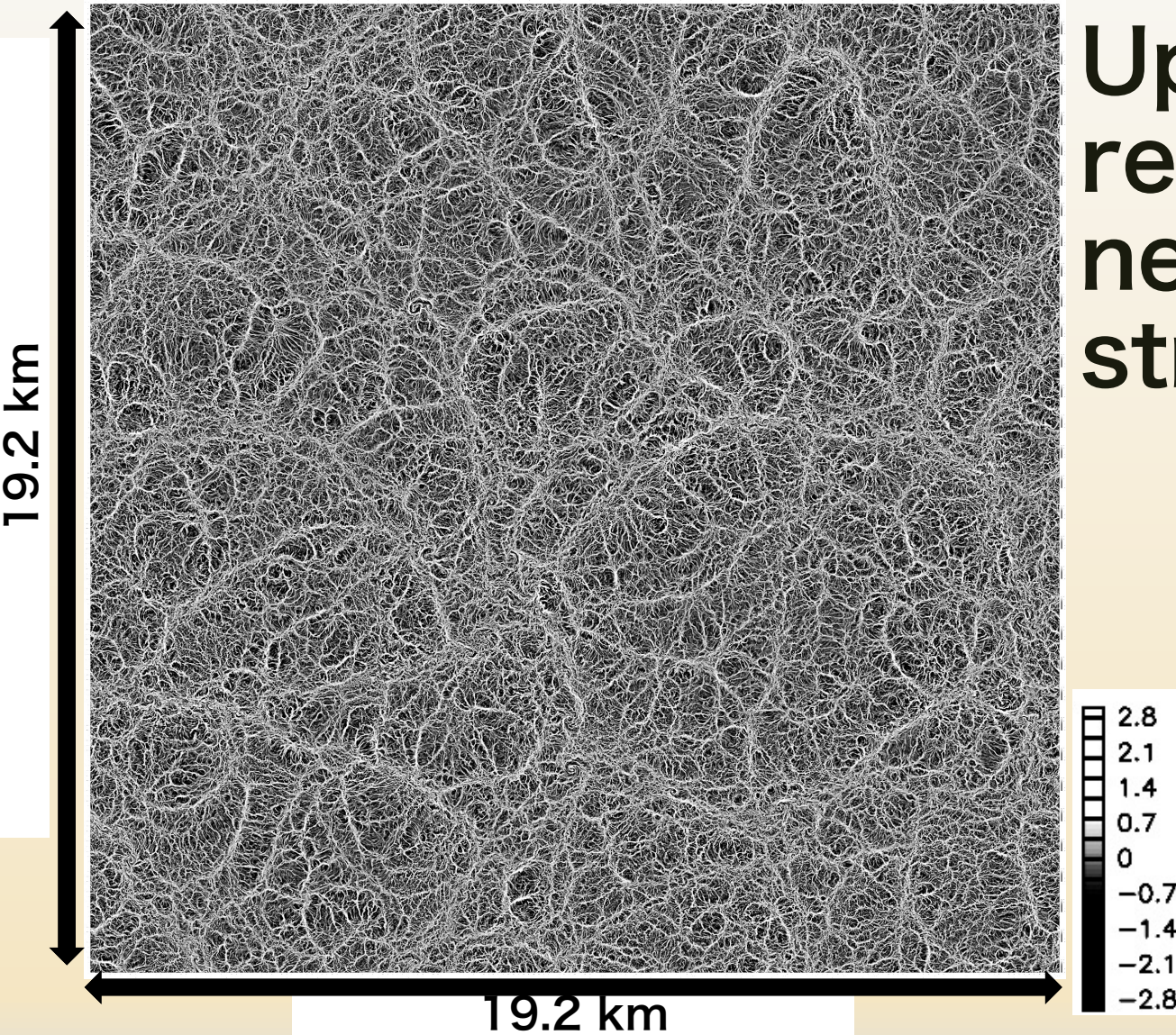
- 10 - 100 m resolution : from LT = 0:00 - 24:00
(LT : Local Time)
- 5 m resolution : from LT = 14:00 - 15:00

Analysis in this study

- In this analysis, using LT = 14 : 30 data.
 - The same as Nishizawa et al. (2016) analysis.
- Analysis procedure
 - Merging original data sets.
 - 5 m resolution original data files consist 7,200 files.
 - It takes a few hour.
 - Each data are about 240 MB.
 - Script language Ruby is used.
 - Making
 - Horizontal distribution
 - Vertical distribution
 - Histogram ...etc.

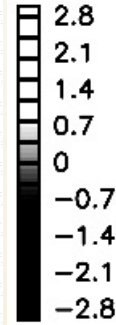
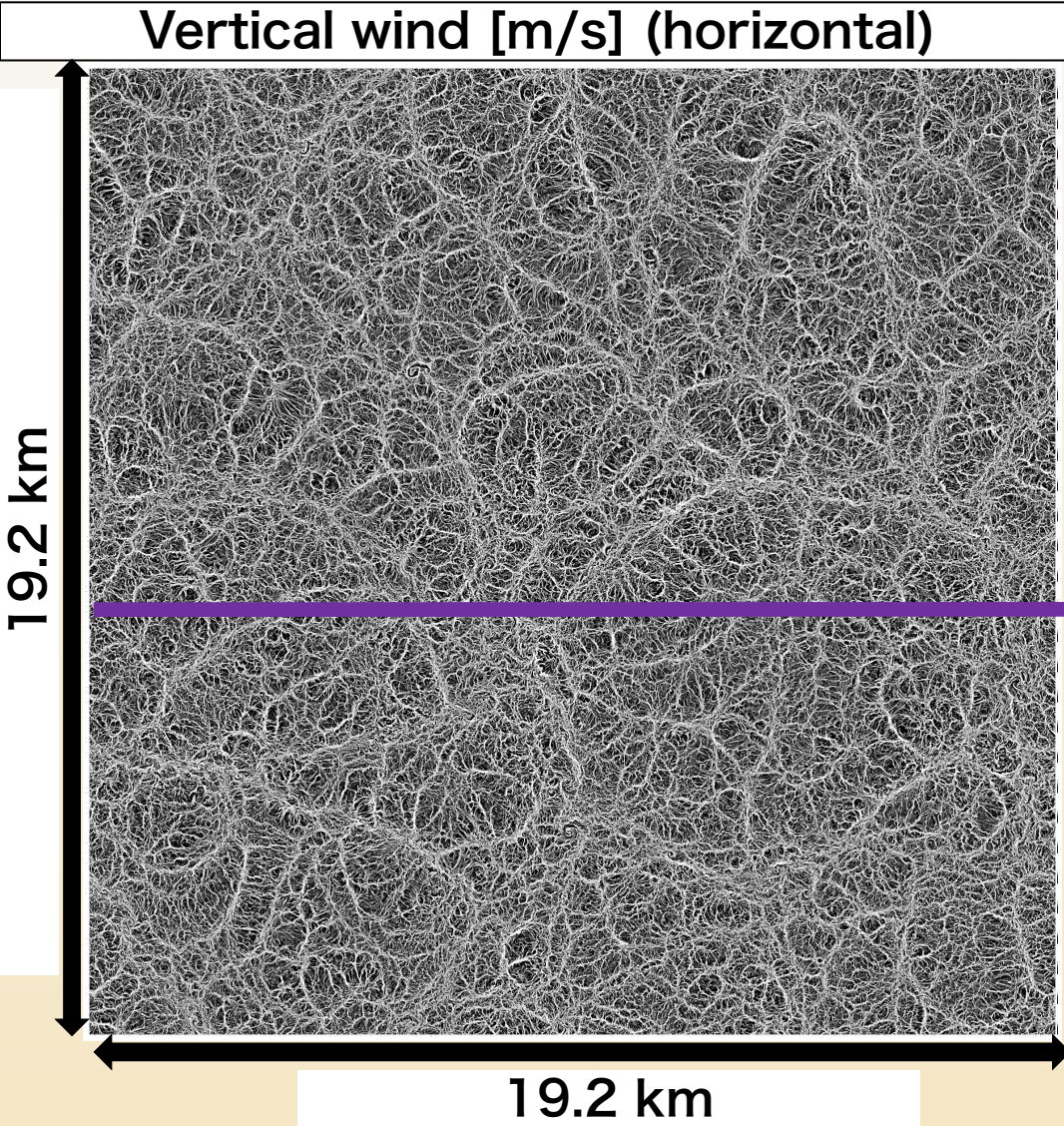
Vertical wind (bottom level $z = 2.5$ m)

Vertical wind [m/s] (horizontal)

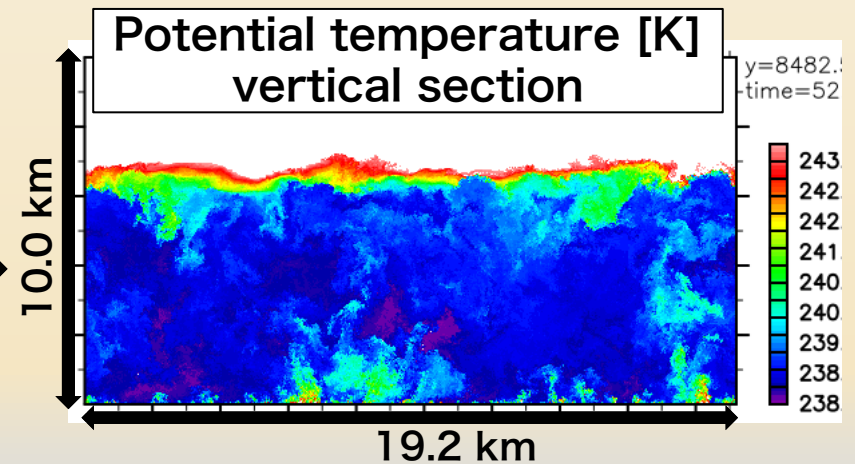
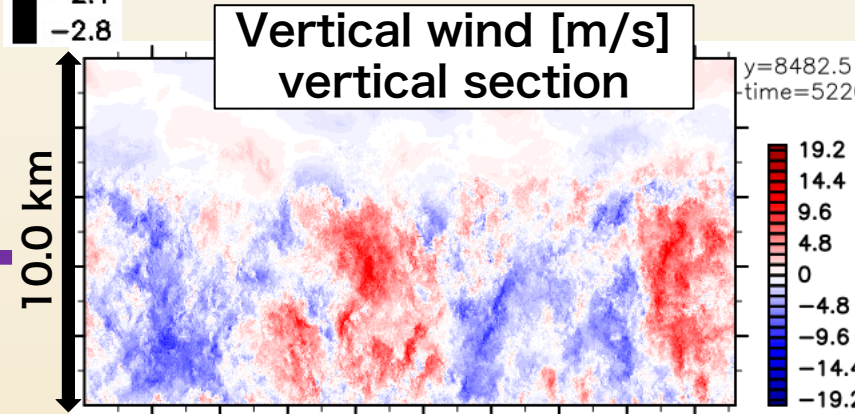


Upward wind region forms network-like structures.

Convective cellular structures



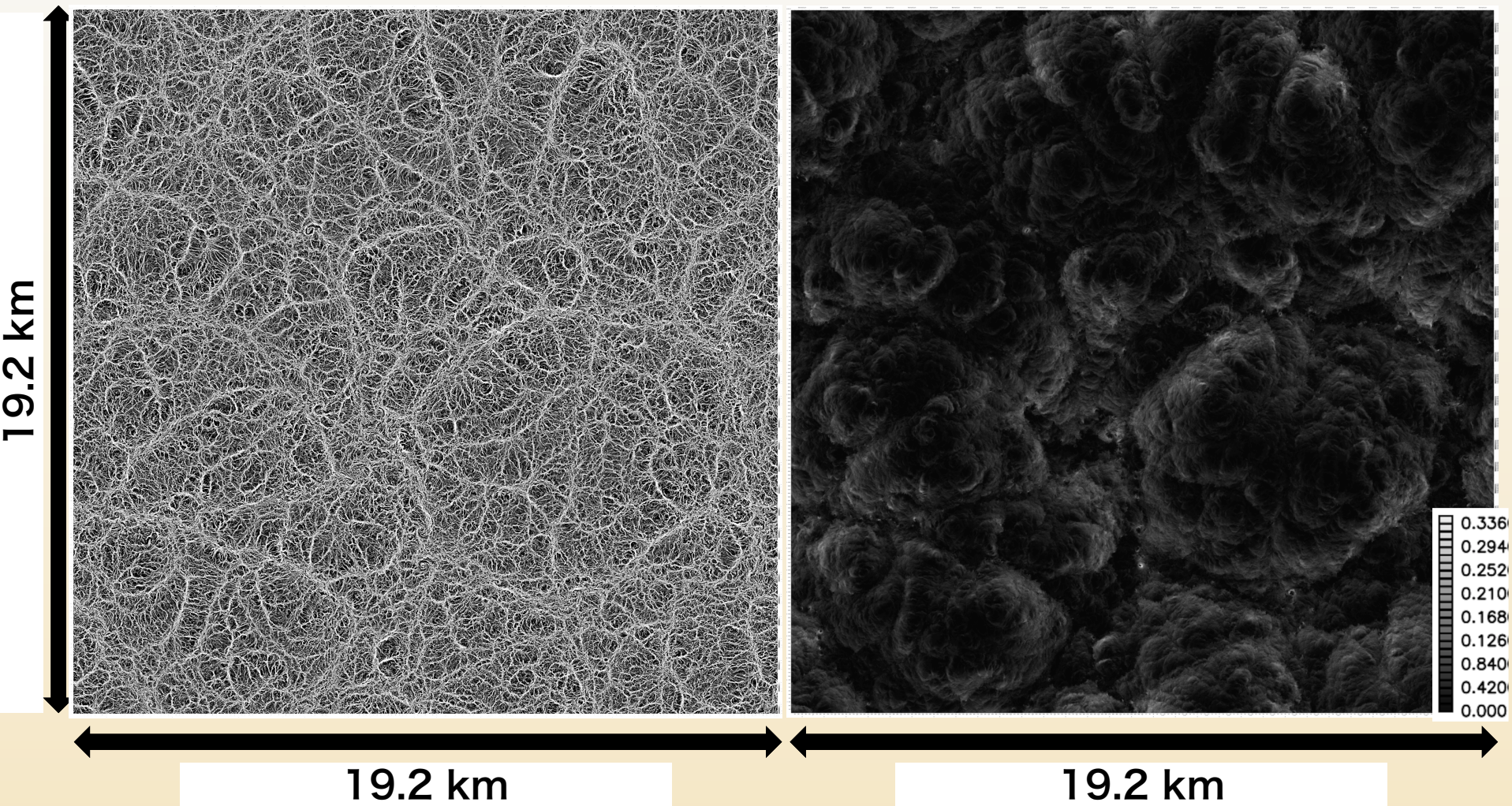
- Convective cell
 - Horizontal scale : Several kilometers
 - Vertical scale : about 6 km



Surface wind stress

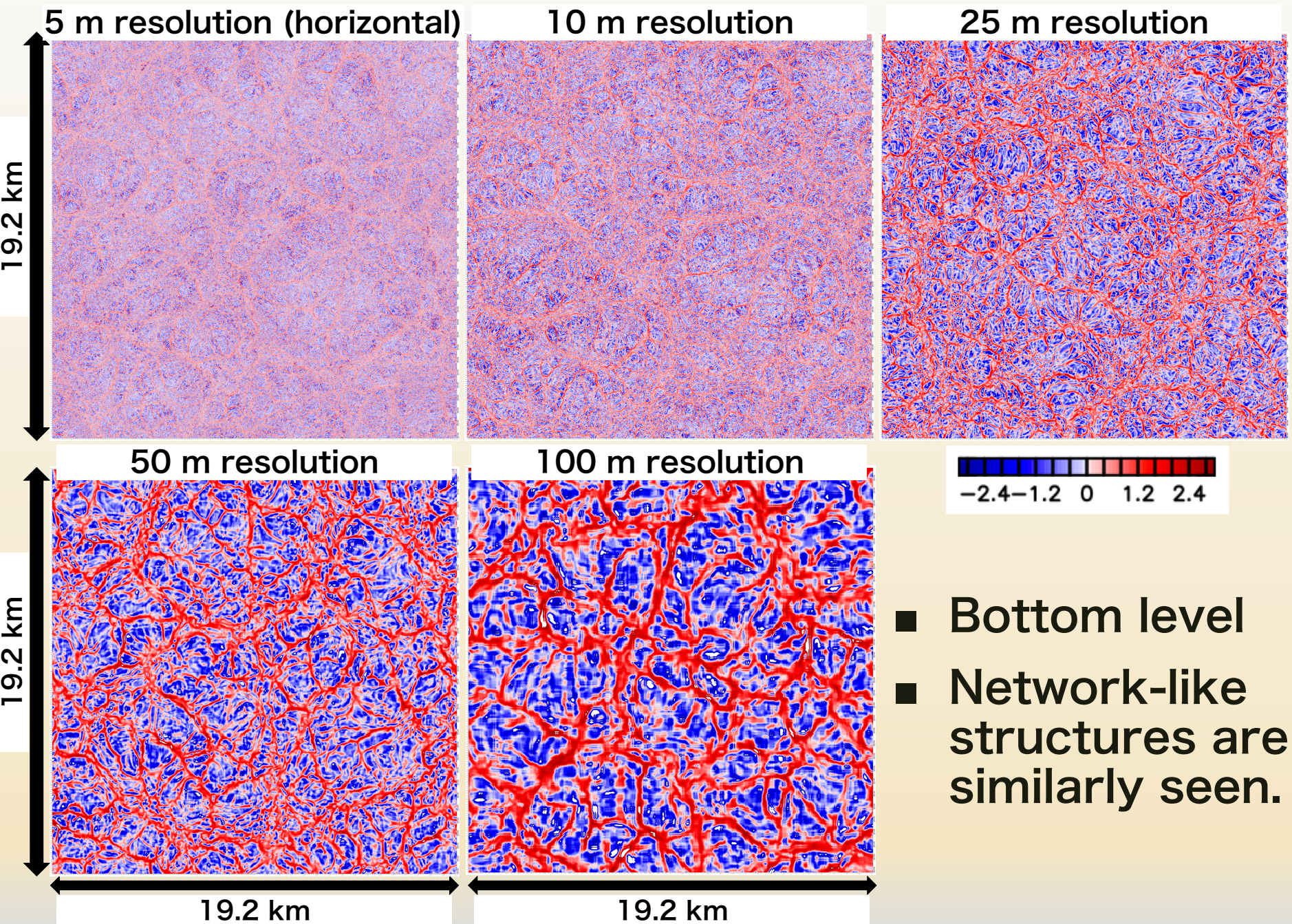
Vertical wind [m/s] (horizontal)

Surface wind stress [Pa] (horizontal)

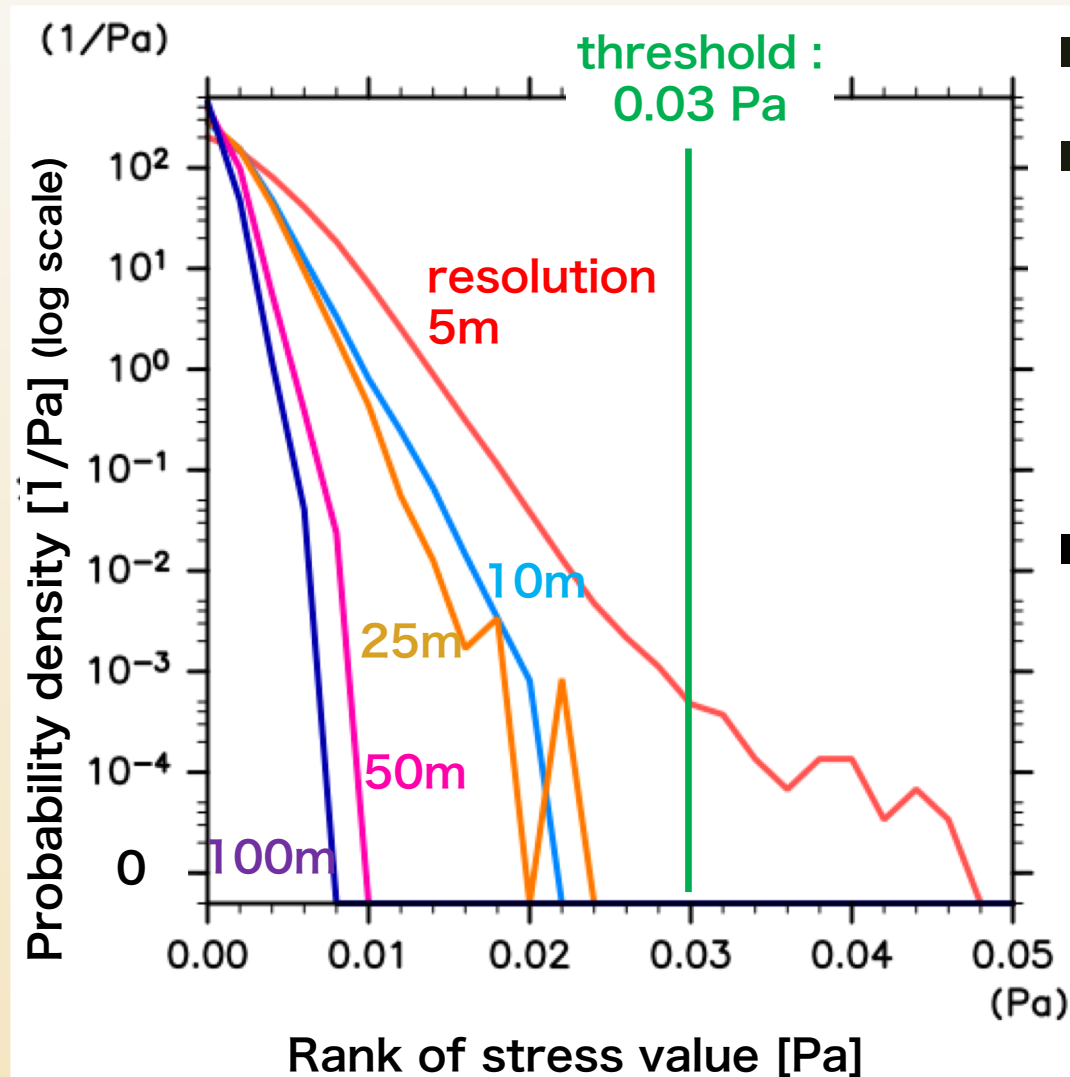


Wind stress tends to be strong near the convective cell boundary

Vertical wind fields of each simulation



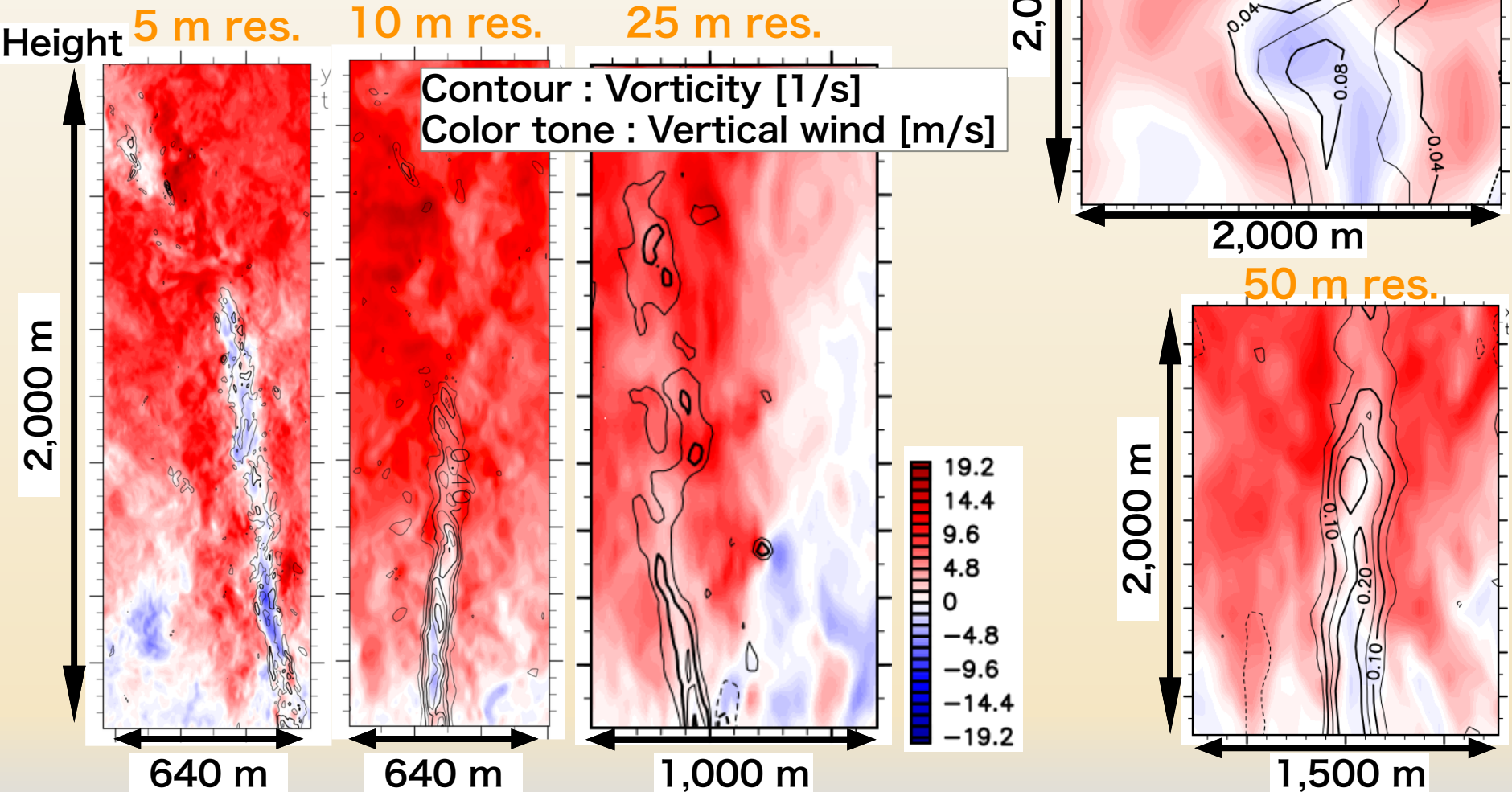
Surface stress probability density distribution



- Bin width : 0.002 Pa
- Result of 5 m resolution is greatly differs from those of more than 10 m resolution results.
- Only 5 m resolution result has the points exceeding threshold value.
- Threshold value 0.03 Pa obtained by experimental results.
(Greeley and Iversen, 1985)

Wind structure associated with the strongest wind stress : vertical sections

- Isolated vortex taller than 1,000 m exists in each resolution.

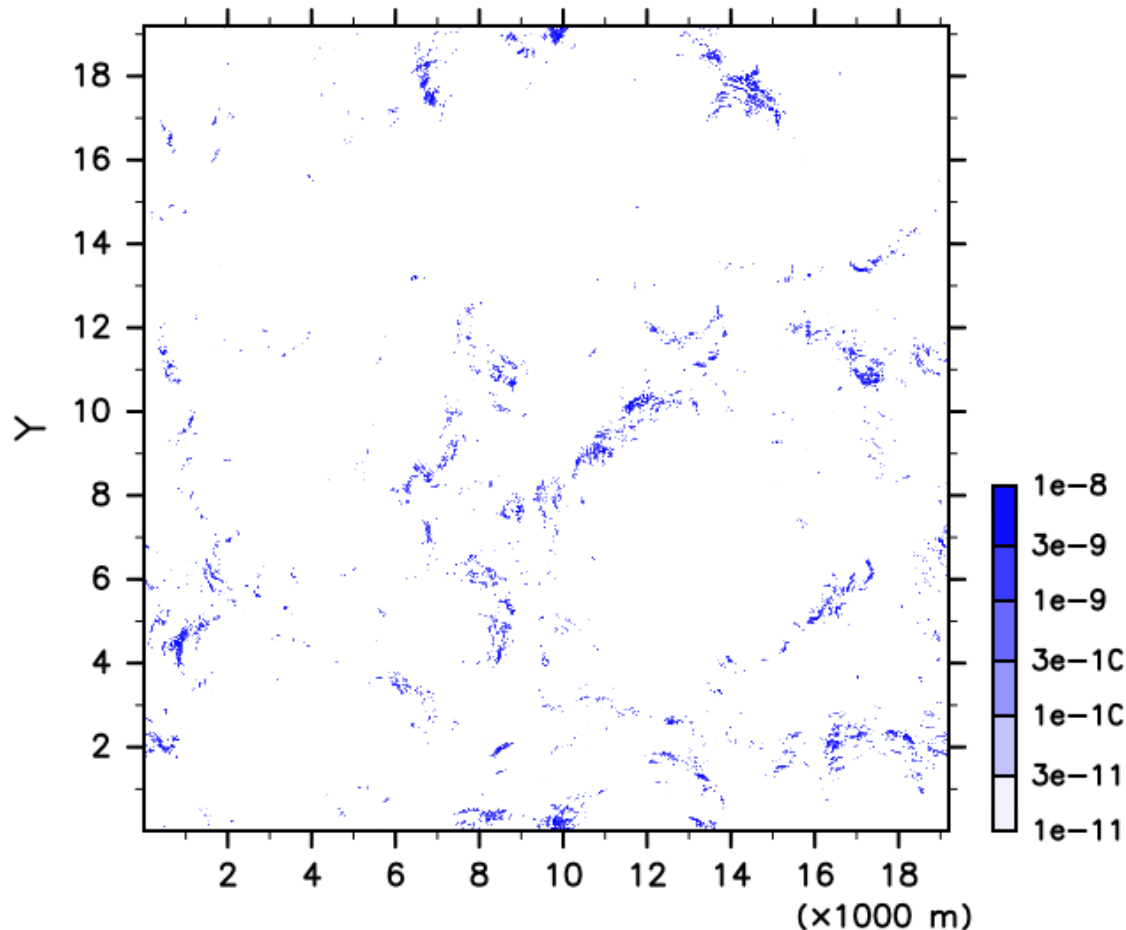


Horizontal distribution of dust flux

- Dust flux is calculated using wind stress scheme (KMH scheme) of Kahre et al. (2006)

Dust flux [kg/m²/s] (horizontal)

(x1000 m)



Horizontal averaged dust flux
[10⁻¹³ kg/(m² s)]

Resolution	5 m	10 m	25 m
Dust flux	391	2.89	3.64

■ Adopted parameters

$$\tau^* : 0.001 \text{ Pa}$$

$$\alpha_W : 0.02$$

- 5 m result has extraordinarily large dust flux

Summary

- Dust lifting parameterization schemes in MGCM have problems.
 - Schemes have been developed without considering details of wind structures.
- Our purpose is to reconsider schemes with examining wind structures.
- We are investigating high-resolution LES for considering validity of parameterization.
 - An isolated vortex taller than 1,000 m exists at the point with the strongest wind stress in each resolution result.
 - Only 5 m resolution result has the points exceeding threshold value of surface wind stress.

References

- Greeley, R., and J. D. Iversen, 1985: Wind as a Geological Process on Earth, Mars, Venus, and Titan., Cambridge Univ. Press., 333 pp
- Kahre, M. K., et al., 2006: Modeling the Martian dust cycle and surface dustreservoirs with the NASA Ames general circulation model, J.G.R., 111, 25
- Louis, J.-F., 1979: A parametric model of vertical eddy fluxes in the atmosphere, Boundary Layer Meteorol., 17, 187-202.
- Mulholland, D. P., et al., 2013: Simulating the interannual variability of major dust storms on Mars using variable lifting thresholds, Icarus, 223, 344-358
- Nishizawa, S., et al., 2016: Martian dust devil statistics from high-resolution large-eddy simulations, Geophys. Res. Lett., 43, 4180-4188
- Odaka M., 2001: A numerical simulation of Martian atmospheric convection with a two-dimensional anelastic model: A case of dust-free Mars, Geophys. Res. Lett., 28, 895-898
- Rennò, N. O., et al., 1998: A simple thermodynamical theory for dust devils, A.M.S., 55, 3244-3252
- Smith, M. D., 2009: THEMIS observations of Mars aerosol optical depth from 2002-2008, Icarus, 202, 444-452
- Westphal, D. L., et al., 1987: A two-dimensional numerical investigation of the dynamics and microphysics of saharan dust storms, J.G.R., 92, 3027-3049
- Wilson, R. J., and Hamilton, K., 1996: Comprehensive model simulation of thermal tides in the Martian atmosphere, J.A.S, 53, 9, 1290-1326

